



1       **First incoherent scatter measurements and adaptive suppression of**  
2 **field-aligned irregularities by the PANSY radar at Syowa Station, Antarctic**

3 Taishi Hashimoto<sup>\*</sup>, Akinori Saito<sup>†</sup>, Koji Nishimura<sup>‡</sup>, Masaki Tsutsumi<sup>§</sup>, Kaoru Sato<sup>¶</sup>, and Toru  
4 Sato<sup>||</sup>

5                   *National Institute of Polar Research, Tachikawa, Japan*

6 <sup>\*</sup>*Corresponding author address:* National Institute of Polar Research, 10-3, Tachikawa, Tokyo,  
7 190-8518, Japan

8 E-mail: hashimoto.taishi.45z@kyoto-u.jp

9 <sup>†</sup>Division of Earth and Planetary Sciences, Kyoto University, Kyoto, Japan

10 <sup>‡</sup>National Institute of Polar Research, Tachikawa, Japan

11 <sup>§</sup>National Institute of Polar Research, Tachikawa, Japan

12 <sup>¶</sup>Department of Earth and Planetary Science, The University of Tokyo, Tokyo, Japan

13 <sup>||</sup>Department of Communications and Computer Engineering, Kyoto University, Kyoto, Japan

**Early Online Release:** This preliminary version has been accepted for publication in *Journal of Atmospheric and Oceanic Technology*, may be fully cited, and has been assigned DOI 10.1175/JTECH-D-18-0175.1. The final typeset copyedited article will replace the EOR at the above DOI when it is published.

## ABSTRACT

14 The Program of the Antarctic Syowa Mesosphere-Stratosphere-  
15 Troposphere/Incoherent Scatter (PANSY) radar is a large atmospheric radar  
16 located at the Antarctic Syowa Station (69.01°S, 39.59°E). The PANSY radar  
17 performed the first incoherent scatter (IS) measurements in the Antarctic re-  
18 gion in 2015. Several specific observations were undertaken in 2017 includ-  
19 ing a 24-h observation of the ionosphere using a peripheral antenna array to  
20 suppress interference from the field-aligned irregularities (FAIs). This paper  
21 presents the preliminary results derived from the IS measurements using the  
22 PANSY radar and the adaptive signal processing techniques to suppress FAIs.  
23 The norm-constrained and directionally-constrained minimization of power  
24 (NC-DCMP) algorithm was applied to the 24-h ionosphere observations by  
25 the PANSY radar with a weighting applied to the directional constraint based  
26 on the gain differences of the subarrays. Compared with the conventional  
27 nonadaptive approach, the number of usable power profiles was increased by  
28 about 24 % by the gain-weighted NC-DCMP algorithm, suggesting its effec-  
29 tiveness for FAI clutter suppression in ionosphere observations. Furthermore,  
30 detection of FAIs using the dedicated antenna array was found valuable in as-  
31 sessing the reliability of estimations of electron density based on VHF-band  
32 IS radar data.

## 33 **1. Introduction**

34 Incoherent scatter (IS) radar is currently the most powerful tool available to investigate the iono-  
35 sphere because it covers a wide altitudinal range from 100 km to 1000 km, and it observes essen-  
36 tial ionospheric parameters such as electron density, ion velocity, ion and electron temperatures,  
37 as well as ion compositions (Evans 1969). The number of the IS radars is limited because they  
38 have high power demand and require a wide antenna aperture. The technical challenges related to  
39 using IS radar have been ongoing for more than five decades. The earliest IS radars, such as those  
40 at Millstone Hill (MA, USA), Jicamarca (Peru) and Arecibo (Puerto Rico) have been operating  
41 since the 1960s, whereas the radars of the European Incoherent Scatter Facility (EISCAT) and the  
42 middle and upper atmosphere (MU) radar at the Shigaraki MU Observatory in Japan began mak-  
43 ing observations in the 1980s. More recently, in the 2000s, Advanced Modular Incoherent Scatter  
44 Radars (AMISR) have been deployed at Poker Flat (AK, USA) and Resolute Bay (Canada). Cur-  
45 rently, the new EISCAT radar, EISCAT\_3D, is in its development phase but it should be operational  
46 by the 2020s (McCrea et al. 2015).

47 All the previously listed IS radars are in the Northern Hemisphere. In contrast, the Program of  
48 the Antarctic Syowa Mesosphere-Stratosphere-Troposphere/Incoherent Scatter (PANSY) radar is  
49 a large atmospheric radar in the Antarctic region, located at Syowa Station (69.01°S, 39.59°E).  
50 As its name states, the PANSY radar has the capability of an IS radar. The PANSY radar per-  
51 formed the first IS radar observations in the Southern Hemisphere in 2015. Clearly, observations  
52 in the Southern Hemisphere are crucial to revealing global features of both the atmosphere and the  
53 ionosphere. In fact, the coupling between the lower and the upper atmospheric layers is expected  
54 to show a large difference between the hemispheres, given the strong hemispheric asymmetry of  
55 the lower atmosphere. Recent in situ observations by the Gravity Field and Steady-State Ocean

56 Circulation Explorer (GOCE) satellite have revealed that atmospheric gravity-wave activity at al-  
57 titudes around 250 km shows marked hemispheric differences. These differences are attributed to  
58 wave activities within the lower atmosphere and the effects of geospatial topology (Trinh et al.  
59 2018). Atmospheric gravity waves in the ionosphere have been observed as traveling ionospheric  
60 disturbances (TIDs) (Hines 1960), which have been investigated using electron density observa-  
61 tions of IS radars at high latitudes (Crowley et al. 1984; Nicolls and Heinselman 2007; Medvedev  
62 et al. 2015) and mid-latitudes (Thome 1964; Oliver et al. 1994) of the Northern Hemisphere. The  
63 PANSY radar observation of the electron density of the Antarctic ionosphere is expected to shed  
64 new light on the coupling between the lower and the upper atmosphere, as well as providing infor-  
65 mation on neutral and ionized components through investigation of TIDs.

66 Although the initial results of the PANSY radar were mainly free from contamination, Sato  
67 et al. (2014) highlighted that strong coherent echoes called ionospheric field-aligned irregularities  
68 (FAIs) can cause severe interference in observations of incoherent scattering of VHF-band radar.  
69 As Fukao et al. (1988) first reported, the MU radar, which uses frequencies similar to the PANSY  
70 radar, also observes FAIs because 3-m-scale (or half wavelength) FAIs are predominant. FAIs are  
71 irregular structures of high electron density aligned along the geomagnetic field lines generated by  
72 plasma instabilities. Radio waves perpendicular to the FAIs are coherently backscattered. Such  
73 scattering conditions are satisfied at around 100 km in height and 30° elevation angles at Syowa  
74 Station when using the VHF band (Ogawa 1997; Koustov et al. 2001). Therefore, the distance of  
75 these FAIs to the radar is greater than 250 km, which causes them to have the same signal range as  
76 the IS in the F region observed by the main array at 70° elevation. To suppress interference from  
77 these FAIs, the PANSY radar has two special antenna arrays used for adaptive signal processing  
78 in cooperation with the main array. These peripheral antenna arrays (hereafter, the FAI array)  
79 were not available during the first period of observation in 2015, but they began operating in 2017.

80 Thus, in 2017, several IS measurements were undertaken using the FAI array, including the 24-h  
81 ionosphere observation described herein.

82 In this paper, we present preliminary results of the ionosphere observations using the PANSY  
83 radar and we describe the methodologies used to suppress interference from the FAIs. Initially,  
84 we review the specifications of the PANSY radar and its supplemental antenna arrays for detecting  
85 FAIs. This is followed by a description of two observations made without and with these arrays in  
86 2015 and 2017, respectively. Next, we outline an adaptive signal processing procedure, optimized  
87 for the current radar system, and we describe the procedure that has been applied to undertake  
88 these observations. We show the results of the application of adaptive signal processing to a 24-h  
89 observation by the PANSY radar in 2017, and we discuss its FAI suppression capabilities. Finally,  
90 we present our concluding remarks concerning the processing of IS radar data with the VHF band.

## 91 **2. Observations**

92 In this section, the specifications of the PANSY radar system and the details of the two observa-  
93 tions made using this radar in 2015 and 2017 are described.

### 94 *a. Receiver array configuration*

95 The main array of the PANSY radar has 1045 three-element crossed-Yagi antennas arranged in  
96 a distributed manner, as shown in Fig. 1 (Sato et al. 2014). In ionosphere observations, all output  
97 signals from the antennas in the main array are combined into a single channel. In addition to this  
98 main array, the PANSY radar has the FAI array, which comprises a pair of peripheral linear arrays  
99 of 12 three-element Yagi antennas configured to observe FAI echoes. Signals from sets of three  
100 adjacent antennas in the FAI array are combined in phase, constituting the eight-channel subarray

101 shown as rectangles 2–9 in Fig. 1. Therefore, this configuration can use adaptive signal processing  
102 techniques for nine subarrays.

103 The FAIs are observed in directions where the magnetic field and the line of sight cross perpen-  
104 dicularly. These conditions are satisfied only at the elevation angle of about  $30^\circ$  at Syowa Station.  
105 Hence, the antennas of the FAI array are directed to the southeast for channels 2–5, and to the  
106 south for channels 6–9, each with an elevation angle of  $30^\circ$ .

107 The antenna patterns of the main and FAI arrays in a section having an azimuth angle of  $135^\circ$   
108 measured clockwise from north are shown in Fig. 2. Because the number of combined antennas  
109 comprising the main and FAI arrays is different, the gain difference to the zenith is about 40 dB  
110 (Fig. 2).

#### 111 *b. First incoherent scatter measurement in 2015*

112 The first IS measurement using the PANSY radar was made on 15 May 2015 during 11:32 –  
113 12:43 local time (LT=UT+3). The methods adopted for obtaining both the observation and the  
114 estimation of the electron density were based on Sato et al. (1989). The parameters for the radar  
115 system are listed in Table 1. During this observation period, signals received from all antennas in  
116 the main array were combined in phase. At this time, the FAI array had not been yet installed and  
117 therefore further signal processing could not be applied.

118 The observed height profile of the electron density averaged over the entire period (71 min)  
119 and all the beam directions is shown in Fig. 3. The coefficient for converting the echo power  
120 into electron density was determined using the  $f_oF_2$  value measured by the on-site ionosonde.  
121 The dashed line in this figure shows the detectability threshold ( $T_d = 3$ ), which is generally used  
122 in atmospheric radars to distinguish a signal from noise (Fukao et al. 2014). As shown by the

123 detectability threshold, the upper limit for IS observation by the PANSY radar is around 700 km.  
124 The noise level is estimated using the region above this height.

125 In this case, the observed height profile followed a reasonable shape for the distribution of the  
126 electron density, implying that this measurement was not affected by the FAIs. Although the  
127 occurrence of E-region FAIs could not be determined by the on-site ionosonde, this assumption is  
128 partially supported by its result for the F region, since no spread F events were recorded during  
129 this period. From these observations, the geomagnetic activity measured using the  $K_p$ -index was  
130 considered low, i.e., it had a  $K_p$  value of 2.

### 131 *c. 24-h continuous observation in 2017*

132 Following the success of the first measurement, a continuous observation of IS was performed  
133 from 13:12 LT on 18 December 2017 until 15:03 LT on 19 December 2017. The observation  
134 parameters were the same as in 2015 and as listed in Table 1. During the 2017 observation period,  
135 the main array was combined in phase, as was done for the 2015 observation. In addition, the FAI  
136 array was enabled, although one of the channels in the FAI array ('9' in Fig. 1) was not working  
137 because of a system malfunction. Thus, seven channels in the FAI array were available to suppress  
138 interference from the FAIs, which yielded eight subarrays.

139 Figure 4 shows an example of height profiles of electron density measured during 13:12 – 15:03  
140 LT on 18 December 2017, averaged over all beam directions. The black line shows the output only  
141 from the main array and the red line is the signal after adaptive signal processing, as explained in  
142 section 3. Other components of the figure are the same as Fig. 3. As described in more detail in  
143 section 4, this observation suffered severe interference from FAIs, in contrast to the observation of  
144 2015. Hence, suppression of the FAIs to estimate the true power of the IS using adaptive signal  
145 processing techniques was required in this case.

146 As shown in Fig. 4, the interference of the FAIs can be seen mainly from 700 km to 1000 km  
147 for this period and adaptive signal processing mitigated these FAIs. For example, around 800 km,  
148 the red line (after adaptive signal processing) becomes below or close to the dashed line (the  
149 detectability threshold), indicating that the contamination from FAIs was successfully suppressed  
150 to reach the noise floor level.

151 Of course, however, the primary concern of this paper is the performance below this region, and  
152 we discuss it in the detailed comparison in section 4. From that perspective, this figure also shows  
153 the stability of the technique when FAI interference is weak because the signal after adaptive signal  
154 processing (red line) is not changed markedly in comparison with the nonadaptive output (black  
155 line) in this region.

### 156 **3. Signal processing**

157 In this section, the basic methodologies involved in adaptive beamforming and their implemen-  
158 tation during the observation are explained.

#### 159 *a. Signal processing algorithm*

160 In 2017, continuous observation was performed using the FAI array. Hence, the output sig-  
161 nals from all the subarray channels were processed using an adaptive beamforming technique to  
162 suppress the FAI echoes.

163 The methodology described herein is based on the directionally constrained minimization of  
164 power (DCMP) algorithm, with an additional constraint on the norm of the weight vector intended  
165 to limit the increase in noise to a designated value. This is known as the norm-constrained DCMP  
166 (NC-DCMP) algorithm (Kamio et al. 2004; Nishimura et al. 2012). Furthermore, the gain differ-



167 ence between the main and FAI arrays must be considered to ensure the norm constraint works  
 168 properly (Hashimoto et al. 2016). The procedures are described below.

169 1) NC-DCMP ALGORITHM

170 The synthesized output of the NC-DCMP algorithm,  $Y$ , is calculated using the following set of  
 171 equations:

$$Y = \mathbf{W}^H \mathbf{X}, \quad (1)$$

$$\underset{\mathbf{W}}{\text{minimize}} \mathbf{W}^H \mathbf{R} \mathbf{W} \text{ subject to } \mathbf{C}^H \mathbf{W} = 1, \quad (2)$$

$$\|\mathbf{W}\|^2 \leq U, \quad (3)$$

172 where  $\mathbf{X} = [X_1, X_2, \dots, X_M]^T$  is a complex time series received by  $M(= 8)$  spatially distributed  
 173 receivers,  $\mathbf{W}$  is the weight vector,  $\mathbf{R} = E[\mathbf{X}\mathbf{X}^H]$  is the covariance matrix,  $\mathbf{C}$  is the directional  
 174 constraint, and  $U$  is the norm constraint. Here,  $E[(\cdot)]$  denotes the ensemble average and  $\|(\cdot)\|$   
 175 represents the Euclidean norm. The norm constraint  $U$  is calculated from the designated noise  
 176 level increase in dB  $L_{\text{dB}}$ :  $U = 10^{-L_{\text{dB}}/10}$ .

177 There is a trade-off between the noise level increase and the clutter suppression capability. How-  
 178 ever,  $L_{\text{dB}} = 0.5$  dB is generally used, because it provides reasonable clutter suppression at the cost  
 179 of small increase in the noise level (Hashimoto et al. 2016). It should also be noted that  $L_{\text{dB}}$  speci-  
 180 fies the worst case, and the actual noise level increase would be smaller if the interference was not  
 181 severe. Because the algorithm minimizes the total output power including noise and interference,  
 182 it automatically suppresses the noise level increase in such cases.

183 2) GAIN WEIGHTING

184 The directional constraint  $\mathbf{C}$  in Eq. (2) is weighted by the power directional gain of the subarrays

185  $\mathbf{G} = [G_1(\theta_o, \phi_o), G_2(\theta_o, \phi_o), \dots, G_M(\theta_o, \phi_o)]^T$  to the desired direction (zenith, azimuth) =  $(\theta_0, \phi_0)$ :

$$C_i = \sqrt{\frac{MG_i(\theta_o, \phi_o)}{\sum_{i=1}^M G_i(\theta_o, \phi_o)}} A_i(\theta_o, \phi_o), \quad (4)$$

$$A_i(\theta, \phi) = \frac{1}{\sqrt{M}} \exp\left(-j \frac{2\pi}{\lambda} \mathbf{L}_i \cdot \mathbf{V}(\theta, \phi)\right), \quad (5)$$

$$\mathbf{V}(\theta, \phi) = [\sin \theta \sin \phi, \sin \theta \cos \phi, \cos \theta]^T,$$

186 where  $j$  is the imaginary unit,  $\lambda$  is the wavelength,  $\mathbf{L}_i$  is the location of the  $i$ -th receiver,  $\mathbf{A}(\theta, \phi)$  is  
 187 the array manifold vector, and  $\mathbf{V}(\theta, \phi)$  is the radial unit vector to the given direction  $(\theta, \phi)$ . Here,  
 188 the azimuth angle is measured clockwise from north.

189 *b. Application to the 2017 observation*

190 Both the nonadaptive beamforming and the gain-weighted NC-DCMP algorithm were applied  
 191 to the eight-channel nonuniform-gain array of the 2017 observation. Hereafter, the received sig-  
 192 nal from the south-directed beam is used in the analysis, i.e., the desired direction in Eq. (5) is  
 193  $(\theta_0, \phi_0) = (20^\circ, 175^\circ)$ . The main array was already synthesized nonadaptively and assigned to  $X_1$   
 194 in Eq. (1), while the other channels from the FAI array were assigned to  $X_{2...8}$ .

195 For the result from the nonadaptive beamforming, we simply used the signal from the first  
 196 channel,  $X_1$ . Conversely, for the result from the NC-DCMP algorithm, we first calculated the  
 197 optimal weight vector  $\mathbf{W}$  through Eqs. (2) to (5). Then, we substituted it in Eq. (1) to obtain the  
 198 synthesized output with the FAIs suppressed. The parameter set for the NC-DCMP algorithm is  
 199 mentioned below.

200 The gain-weighting coefficients  $\mathbf{G}$  were  $G_1(\theta_0, \phi_0) = 0.9958$  and  $G_{2...8}(\theta_0, \phi_0) = 0.0372$ , deter-  
 201 mined using the power directional gain differences to  $(\theta_0, \phi_0)$ . The norm constraint was set as  $U =$

202 1.122 to limit the increase in the noise to less than 0.5 dB. The covariance matrix  $\mathbf{R}$  was calculated  
 203 using  $N = 256$  samples before and after the target time  $t_i$ , i.e.,  $\mathbf{R}(t_i) = \frac{1}{2N+1} \sum_{j=i-N}^{i+N} \mathbf{X}(t_j) \mathbf{X}^H(t_j)$ .  
 204 Consequently, the duration for averaging the covariance matrix was about 21 s. Note that this  
 205 period is shorter than used for the mesosphere-stratosphere-troposphere region of about 1 min  
 206 (Hashimoto et al. 2016), because the correlation time between the IS and the FAI signal is much  
 207 shorter than between atmospheric echoes and ground clutter.

208 *c. Calculation of the height profiles and meteor rejection*

209 After applying the beamforming technique, height profiles were obtained by averaging the re-  
 210 ceived power. For every  $N_a (= 204)$  samples ( $\sim 8.16$  s), the DC offset was removed and the power  
 211 was averaged for each range to obtain a single height profile  $P(r)$ , in which  $r$  is the height index  
 212 of the profile.

213 To remove contamination from meteors and other artifacts, the following thresholding procedure  
 214 was applied successively. Initially, for every  $N_1 (= 22)$  height profiles ( $\sim 3$  min), the mean noise  
 215 level  $\bar{P}_{N_1}(r)$  was estimated for each range. Then, each profile  $P(r)$  in the set was tested to ascertain  
 216 whether the following condition was satisfied:

$$\left\{ r \mid P(r) \geq \left( 1 + \frac{T_d}{\sqrt{N_a}} \right) \bar{P}_{N_1}(r) \right\} = \emptyset, \quad (6)$$

217 where  $T_d = 7$  is the constant used for thresholding in this study. Only profiles satisfying Eq. (6)  
 218 were used in the incoherent integrations. Here, the left-hand side of Eq. (6) indicates a set of  
 219 height indices where the signal level exceeds the detectability threshold, and  $T_d = 7$  was chosen  
 220 experimentally to balance the impact of the number of discarded profiles and remaining meteor  
 221 echoes on the smoothness of the averaged profiles. Finally,  $N_2 (= 5)$  successive sets of meteor-  
 222 rejected profiles passing the test in Eq. (6) were integrated to calculate 15-min-averaged profiles.

223 However, if the number of averaged profiles was less than 70 % of the total number of profiles  
 224  $N_1 N_2 (= 110)$ , i.e., less than 77 in a 15-min interval, then the entire set of profiles was discarded to  
 225 avoid unreliable estimation. Herein, this procedure is called meteor rejection.

226 *d. Detection of field-aligned irregularities*

227 The procedures for applying the proposed signal processing techniques to the 2017 observation  
 228 have been explained above. For quantitative discussion in section 4 regarding the FAI suppression  
 229 capability of the proposed method, the occurrence of FAIs is first considered.

230 The occurrence of FAIs was determined using the received signals from the FAI array. As shown  
 231 in Fig. 2, the directivity gains of the FAI channels are optimized for directions where FAIs appear,  
 232 with no response in the main beam direction. Therefore, all echoes observed by the FAI channels  
 233 can be considered FAIs or meteors. Here, we did not distinguish between these two echo sources  
 234 because it is advantageous to suppress meteor echoes when they are observed by the FAI channels.

235 For each FAI channel, averaged power profiles without applying the meteor rejection procedure,  
 236  $P^{(i)}(r)$  ( $i = 2, \dots, 8$ ), were first calculated with  $N_1 N_2$  incoherent integrations. Then, the mean noise  
 237 level of the  $i$ -th channel in a 15-min interval,  $\bar{P}_{N_1 N_2}^{(i)}(r)$ , was calculated from the profile  $P^{(i)}(r)$  to  
 238 build the detectability threshold in the same manner as Eq. (6). Hence, any ranges satisfying the  
 239 following condition were assumed to have FAIs:

$$\left\{ r \mid P^{(i)}(r) \geq \left( 1 + \frac{T_d}{\sqrt{N_a N_1 N_2}} \right) \bar{P}_{N_1 N_2}^{(i)}(r) \right\} \neq \emptyset. \quad (7)$$

240 Note that the FAI array is not beamformed, and Eq. (7) is satisfied if any one of the FAI channels  
 241 exceeds the detectability threshold.

## 4. Results and discussions

Figure 5 shows time series of 15-min-averaged power profiles and FAI occurrences from the south-directed beam during the continuous observation period of 2017. The panels of this figure show (a) the raw power profiles, i.e., those obtained from nonadaptive beamforming without meteor rejection, (b) power profiles obtained using the conventional method, i.e., nonadaptive beamforming with meteor rejection, (c) power profiles obtained using the proposed method, i.e., the gain-weighted NC-DCMP algorithm with meteor rejection, and (d) the occurrence of FAIs during the observation period. The horizontal axis of each panel is time (LT) and the vertical axis is range in km multiplied by  $\cos 20^\circ$ , which can be interpreted as actual heights for the results from the main array, i.e., in panels (a) – (c). In the upper three panels, colors indicate the intensity of the estimated power profiles against the noise level in dB. In the bottom panel, the black hatching shows the time and height of FAI occurrence, estimated using the procedure described in section 3.d. Note that these FAIs are considered to originate in E region because the perpendicular condition is only satisfied therein using the VHF band, as mentioned in section 1. This panel also contains data from the on-site ionosonde, illustrating the occurrence of spread F events by the red hatching.

As shown in panels (b) and (c) of Fig. 5, the number of profiles discarded by the meteor rejection process is reduced when the NC-DCMP algorithm is used. The ratio of rejected profiles to total profiles in panel (b) is 35.78 %, while in panel (c) it is 11.92 %, representing an improvement of 23.85 %. The average loss of signals by the NC-DCMP algorithm is about 0.02 dB, which is calculated from the difference between panels (b) and (c) using the times and ranges without the black hatching in panel (d) of Fig. 5, i.e., the data in which the FAIs are not detected. Note that

264 this is sufficiently small value compared with the maximum allowable increase of the noise level  
265 (0.5 dB) designated in section 3.b, indicating the stability of the proposed method.

266 Compared with panel (b), the profiles in panel (c) have less temporal discontinuities in received  
267 power, especially during daytime, e.g., 12:00–15:00 LT on both days. For example, an unnatural  
268 step discontinuity at 12:45 in panel (b) is mitigated, and some missing profiles are restored in  
269 panel (c). In these intervals, panels (b) and (c) differ markedly from panel (a), reflecting the meteor  
270 rejection process, although almost no FAIs or meteors were detected above 250 km, as indicated in  
271 panel (d) of Fig. 5. This is due to a lack of sensitivity in the FAI array, making it difficult to detect  
272 weak echoes above this height using the given detectability threshold. Nevertheless, the proposed  
273 gain-weighted NC-DCMP algorithm did suppress weak clutter to obtain diurnal variation in the  
274 background IS with less rejected profiles and ignorable loss of desired signals in comparison with  
275 existing techniques. This is one of the main advantages of adaptive signal processing.

276 Interestingly, very strong FAIs, e.g., during 19:00–22:00 LT on 18 December 2017, were not  
277 suppressed fully by the adaptive signal processing, as shown in panel (c) of Fig. 5. Even in such  
278 cases, we can still detect FAIs using the FAI array, and we can mark these contaminated records  
279 as requiring caution in their analysis. Hence, this method is more robust in estimating the electron  
280 density than using signals only from the main array. Furthermore, although the on-site ionosonde  
281 detected only the existence of spread F events as shown by the red hatching in panel (d), which  
282 is not necessarily related to the E-region FAIs, the FAI array directly observed the E-region FAIs  
283 and provided a broader coverage in both temporal and spatial distributions. The importance of the  
284 dedicated FAI array is also supported by this result, since the ionosonde and FAI array observe  
285 different targets using different frequencies, i.e., the FAIs in F and E regions using HF and VHF  
286 bands, respectively.

287 From the above discussion, we conclude that the FAI array is effective for assessing the reliabil-  
288 ity of estimations of electron density based on VHF-band IS radar data.

## 289 **5. Summary and conclusions**

290 This paper presents the initial results of ionosphere observations by the PANSY radar, which  
291 uses an adaptive signal processing algorithm for signals from auxiliary antenna arrays to suppress  
292 FAI clutter. The gain-weighted NC-DCMP algorithm applied to the eight-channel subarray con-  
293 figuration of the PANSY radar had an average suppression ratio of 2.17 dB against echoes from  
294 FAIs and meteors. This increased the number of usable power profiles by 23.85 %, compared  
295 with conventional nonadaptive beamforming. In addition, even when strong FAIs could not be  
296 suppressed fully, we can still use the output from the FAI array to detect FAIs and to reject con-  
297 taminated records, which improves the reliability of the estimations of electron density based on  
298 VHF-band IS radar data.

299 From these results, we conclude that using a subarray dedicated to FAI observation, together  
300 with adaptive signal processing, is valuable for ionosphere observations in the Antarctic region.  
301 Furthermore, the gain-weighted NC-DCMP algorithm is acknowledged as a satisfactory workable  
302 solution for the PANSY radar. This proposed method and configuration are planned to be applied  
303 to the other types of IS radar observation of the PANSY radar, e.g., multipulse observations used  
304 to measure ion drift speeds and temperatures, in the future.

305 *Acknowledgments.* PANSY is a multi-institutional project with a core team at both The Univer-  
306 sity of Tokyo and the National Institute of Polar Research. The PANSY radar is operated by the  
307 staff of the Japanese Antarctic Research Expedition. The on-site ionosonde data were provided by  
308 the National Institute of Information and Communications Technology.

## References

- Crowley, G., T. B. Jones, T. R. Robinson, N. M. Wade, and O. Holt, 1984: Determination of the vertical neutral temperature and wind profiles using EISCAT and HF Doppler radar. *Journal of Atmospheric and Terrestrial Physics*, **46**, 501–507, doi:10.1016/0021-9169(84)90068-0.
- Evans, J. V., 1969: Theory and practice of ionosphere study by Thomson scatter radar. *Proceedings of the IEEE*, **57** (4), 496–530, doi:10.1109/PROC.1969.7005.
- Fukao, S., K. Hamazu, and R. J. Doviak, 2014: *Radar for meteorological and atmospheric observations*. Springer, doi:10.1007/978-4-431-54334-3.
- Fukao, S., J. McClure, A. Ito, T. Sato, I. Kimura, T. Tsuda, and S. Kato, 1988: First VHF radar observation of midlatitude F-region field-aligned irregularities. *Geophysical Research Letters*, **15** (8), 768–771.
- Hashimoto, T., K. Nishimura, and T. Sato, 2016: Adaptive sidelobe cancellation technique for atmospheric radars containing arrays with nonuniform gain. *IEICE Transactions on Communications*, doi:10.1587/transcom.2016EBP3047, submitted for publication.
- Hines, C. O., 1960: Internal atmospheric gravity waves at ionospheric heights. *Canadian Journal of Physics*, **38** (11), 1441–1481, doi:10.1029/GM018p0248.
- Kamio, K., K. Nishimura, and T. Sato, 2004: Adaptive sidelobe control for clutter rejection of atmospheric radars. *Annales Geophysicae*, **22** (11), 4005–4012, doi:10.5194/angeo-22-4005-2004.
- Koustov, A. V., K. Igarashi, D. Andr'e, K. Ohtaka, N. Sato, H. Yamagishi, and A. Yukimatu, 2001: Observations of 50- and 12-MHz auroral coherent echoes at the Antarctic Syowa station. *Journal of Geophysical Research: Space Physics*, **106** (A7), 12–875, doi:10.1029/2000JA000165.



- 331 McCrea, I., and Coauthors, 2015: The science case for the EISCAT\_3D radar. *Progress in Earth*  
332 *and Planetary Science*, **2** (1), 21, doi:10.1186/s40645-015-0051-8.
- 333 Medvedev, A., K. Ratovsky, M. Tolstikov, S. Alsatkin, and A. Shcherbakov, 2015: A statistical  
334 study of internal gravity wave characteristics using the combined Irkutsk Incoherent Scatter  
335 Radar and Digisonde data. *Journal of Atmospheric and Solar-Terrestrial Physics*, **132**, 13–21,  
336 doi:10.1016/j.jastp.2015.06.012.
- 337 Nicolls, M. J., and C. J. Heinselman, 2007: Three-dimensional measurements of traveling iono-  
338 spheric disturbances with the Poker Flat Incoherent Scatter Radar. *Geophysical Research Let-*  
339 *ters*, **34** (21), doi:10.1029/2007GL031506.
- 340 Nishimura, K., T. Nakamura, T. Sato, and K. Sato, 2012: Adaptive beamforming technique for  
341 accurate vertical wind measurements with multi-channel MST radar. *Journal of Atmospheric*  
342 *and Oceanic Technology*, **29** (12), 1769–1775, doi:10.1175/JTECH-D-11-00211.1.
- 343 Ogawa, T., 1997: Radar observations of ionospheric irregularities at Syowa Station, Antarctica: A  
344 brief overview. *Annales Geophysicae*, **14** (12), 1454–1461, doi:10.1007/s00585-996-1454-z.
- 345 Oliver, W. L., S. Fukao, Y. Yamamoto, T. Takami, M. D. Yamanaka, M. Yamamoto, T. Nakamura,  
346 and T. Tsuda, 1994: Middle and upper atmosphere radar observations of ionospheric density  
347 gradients produced by gravity wave packets. *Journal of Geophysical Research: Space Physics*,  
348 **99** (A4), 6321–6329, doi:10.1029/94JA00171.
- 349 Sato, K., and Coauthors, 2014: Program of the Antarctic Syowa MST/IS radar (PANSY). *Journal*  
350 *of Atmospheric and Solar-Terrestrial Physics*, **118A**, 2–15, doi:10.1016/j.jastp.2013.08.022.

351 Sato, T., I. Atsuo, W. L. Oliver, S. Fukao, T. Tsuda, S. Kato, and I. Kimura, 1989: Ionospheric  
352 incoherent scatter measurements with the middle and upper atmosphere radar: Techniques and  
353 capability. *Radio Science*, **24** (1), 85–98, doi:10.1029/RS024i001p00085.

354 Thome, G. D., 1964: Incoherent scatter observations of traveling ionospheric disturbances. *Jour-*  
355 *nal of Geophysical Research*, **69** (19), 4047–4049, doi:10.1029/JZ069i019p04047.

356 Trinh, Q. T., M. Ern, E. Doornbos, P. Preusse, and M. Riese, 2018: Satellite observations of  
357 middle atmosphere–thermosphere vertical coupling by gravity waves. *Annales Geophysicae*,  
358 Copernicus GmbH, Vol. 36, 425–444, doi:10.5194/angeo-36-425-2018.

359 **LIST OF TABLES**

360 **Table 1.** Parameters for the incoherent scatter measurements made on 15 May 2015 and  
361 during 18–19 December 2017 by the PANSY radar. . . . . 20

TABLE 1: Parameters for the incoherent scatter measurements made on 15 May 2015 and during 18–19 December 2017 by the PANSY radar.

Center frequency	47 MHz
Ranges	140 km to 1230 km
Range resolution $\Delta r$	9.6 km
Beam zenith $\theta$	20°
Beam azimuth $\phi$	−5°, 85°, 175°, −95°
Time resolution $\Delta t$	40 ms
Pulse compression	7-bit Barker code

362 **LIST OF FIGURES**

363 **Fig. 1.** Antenna positions of the main and FAI arrays of the PANSY radar, and the subarray assign-  
364 ment in the IS observation during 18–19 December 2017. . . . . 22

365 **Fig. 2.** Antenna patterns of the PANSY radar in the section having an azimuth angle of  $135^\circ$  mea-  
366 sured clockwise from north. Solid line (MAIN) is for the main array, dashed line (FAI 1) is  
367 for FAI channels 2–5, and dotted line (FAI 2) is for FAI channels 6–9. . . . . 23

368 **Fig. 3.** First electron density profile observed by the PANSY radar on 15 May 2015. The profile  
369 is averaged about an hour over all beam directions. The dashed line shows the detectability  
370 threshold ( $T_d = 3$ ). The observation parameters are listed in Table 1. The conversion coef-  
371 ficient from echo power to electron density is determined using the  $f_oF_2$  value measured by  
372 an on-site ionosonde. . . . . 24

373 **Fig. 4.** Example electron density profiles observed by the PANSY radar on 18 December 2017. The  
374 profiles are averaged about an hour over all beam directions. Black line shows the output  
375 only from the main array and red line shows the signal after adaptive signal processing, as  
376 explained in section 3. Other components are the same as Fig. 3. . . . . 25

377 **Fig. 5.** Time series of 15-min-averaged power profiles for the south-directed beam during the con-  
378 tinuous observation period of 2017: (a) raw power profiles, (b) power profiles obtained  
379 using nonadaptive beamforming with meteor rejection, (c) power profiles obtained using  
380 the NC-DCMP algorithm with meteor rejection, and (d) occurrence of FAIs estimated using  
381 channels 2–8 (FAI). The bottom panel also contains the occurrence of spread F events de-  
382 termined by the on-site ionosonde (Fs). The horizontal axis of each panel is time (LT) and  
383 the vertical axis is range multiplied by  $\cos 20^\circ$ . . . . . 26

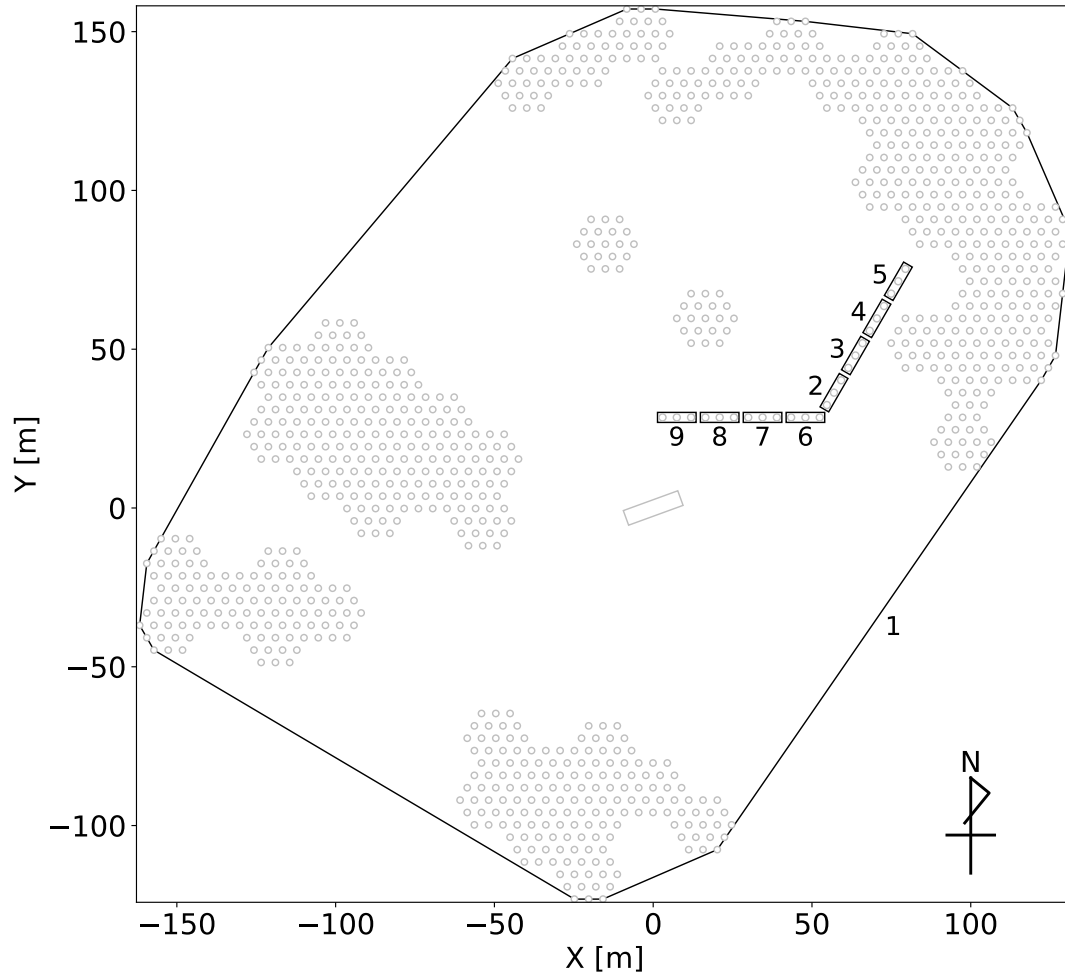


FIG. 1: Antenna positions of the main and FAI arrays of the PANSY radar, and the subarray assignment in the IS observation during 18–19 December 2017.

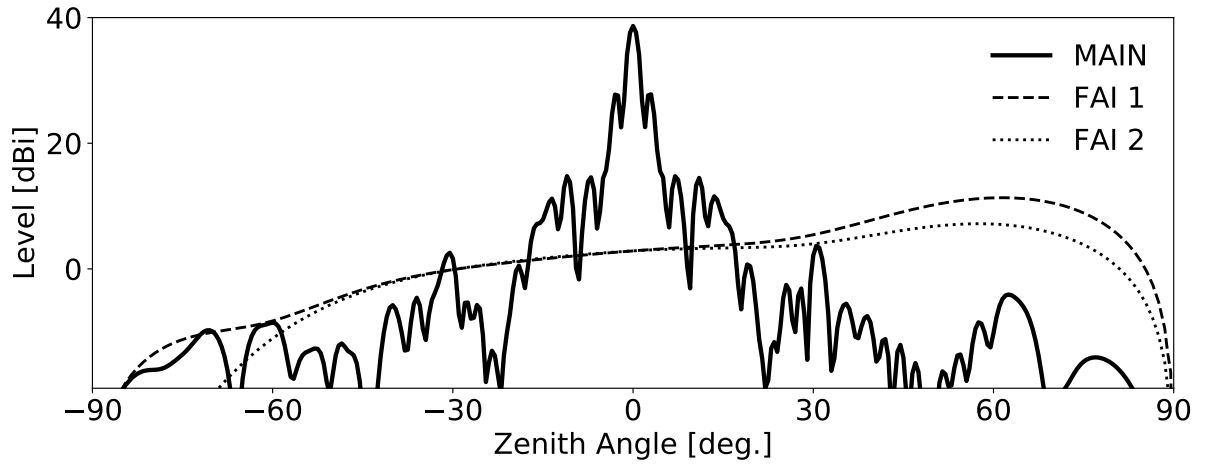


FIG. 2: Antenna patterns of the PANSY radar in the section having an azimuth angle of  $135^\circ$  measured clockwise from north. Solid line (MAIN) is for the main array, dashed line (FAI 1) is for FAI channels 2–5, and dotted line (FAI 2) is for FAI channels 6–9.

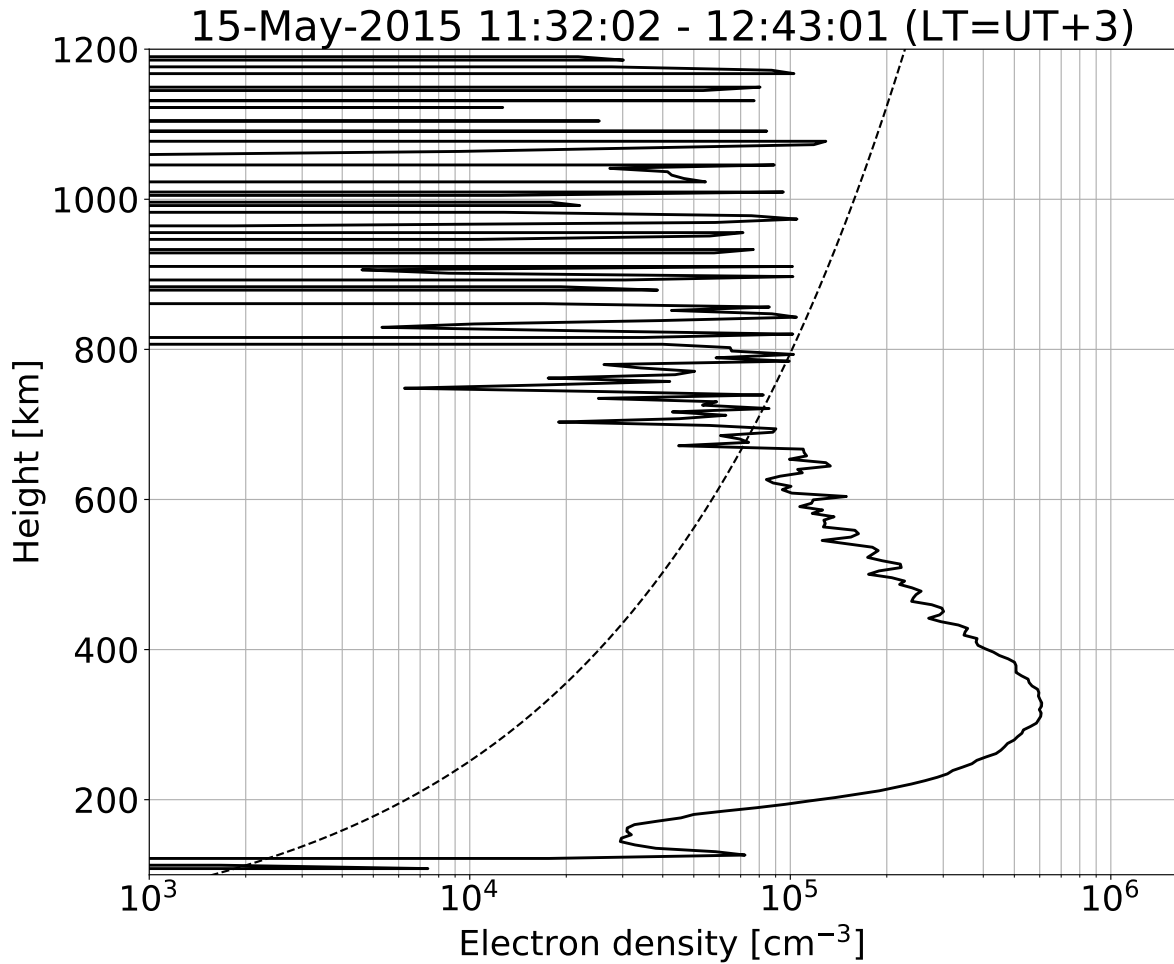


FIG. 3: First electron density profile observed by the PANSY radar on 15 May 2015. The profile is averaged about an hour over all beam directions. The dashed line shows the detectability threshold ( $T_d = 3$ ). The observation parameters are listed in Table 1. The conversion coefficient from echo power to electron density is determined using the  $f_oF_2$  value measured by an on-site ionosonde.



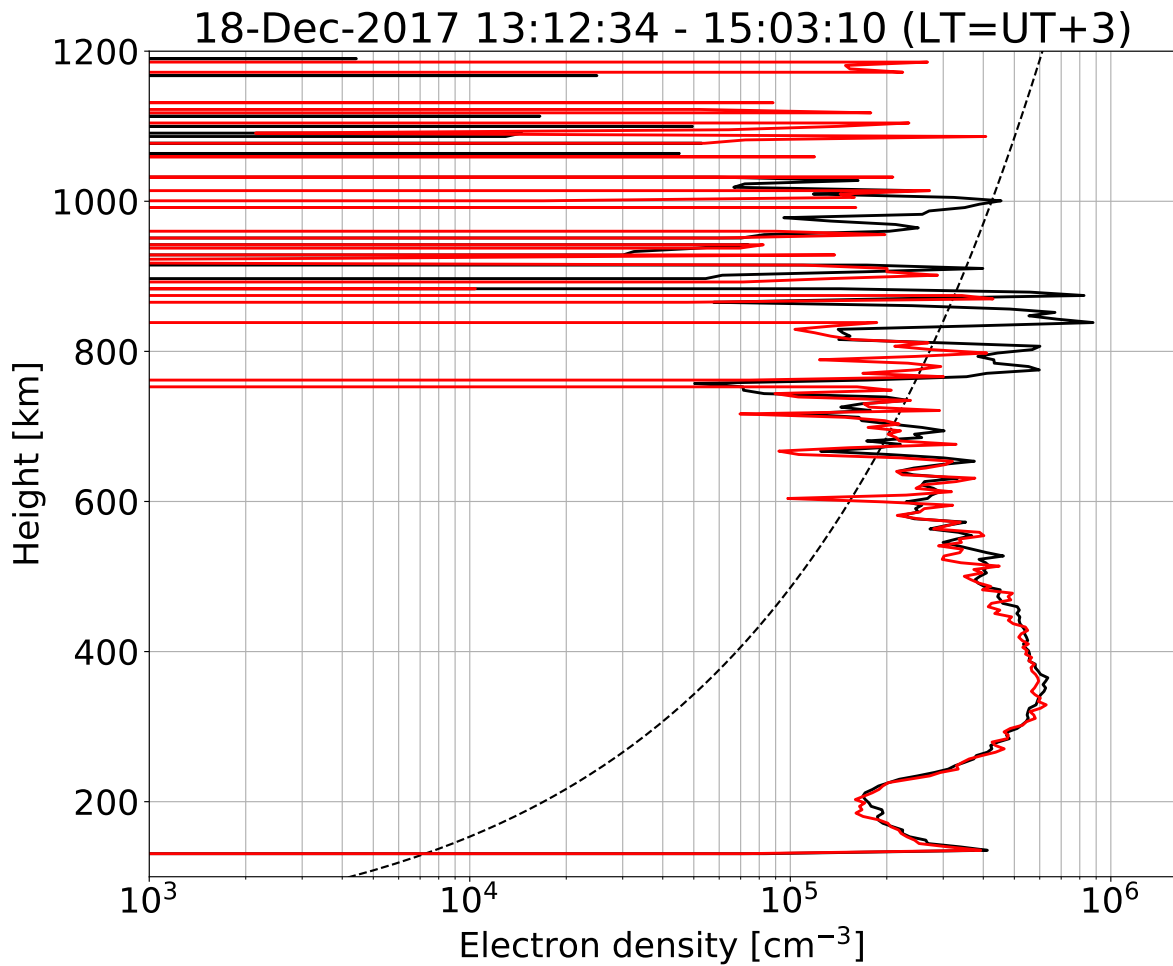


FIG. 4: Example electron density profiles observed by the PANSY radar on 18 December 2017. The profiles are averaged about an hour over all beam directions. Black line shows the output only from the main array and red line shows the signal after adaptive signal processing, as explained in section 3. Other components are the same as Fig. 3.

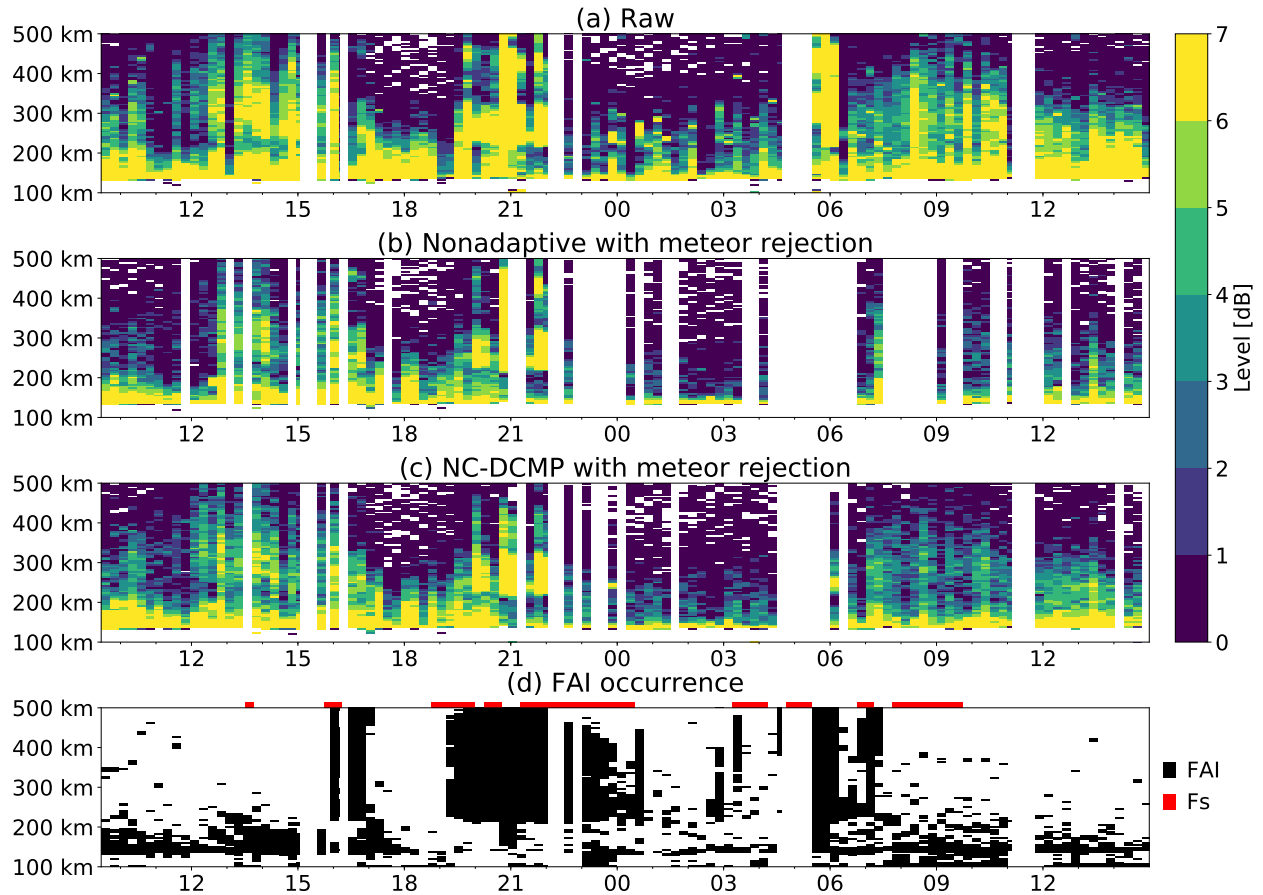


FIG. 5: Time series of 15-min-averaged power profiles for the south-directed beam during the continuous observation period of 2017: (a) raw power profiles, (b) power profiles obtained using nonadaptive beamforming with meteor rejection, (c) power profiles obtained using the NC-DCMP algorithm with meteor rejection, and (d) occurrence of FAIs estimated using channels 2–8 (FAI). The bottom panel also contains the occurrence of spread F events determined by the on-site ionosonde (Fs). The horizontal axis of each panel is time (LT) and the vertical axis is range multiplied by  $\cos 20^\circ$ .